Properties and Characterization of Subaqueous Soils



Martin C. Rabenhorst Environmental Science and Technology University of Maryland



Mark H. Stolt Natural Resources Science University of Rhode Island

What is unique to this system?

Assumption: You are familiar with properties and characterization of soils in general. So what makes subaqueous systems different?

What is unique to this system?

They are permanently submerged

- Impacts sampling big time!
- Affects the properties of the soils themselves
- Air excluded, except from the uppermost zone (mm to cm); strongly anaerobic

They are young (at least in the upper part)

- Mostly late Holocene in age
- Generally weakly developed profiles
- Many have formed in brackish or saline environments
- Sulfate a dominant anion that affects products of reduction

Sampling

- Bucket Auger
 - Limited to shallow water must sample in water
 - Fast and inexpensive
 - Poor horizon resolution
 - May have difficulties with high n-value material
 - No volume controls (for Bulk Density)
 - During warm season or dry suits needed



Sampling

- MacCauley Sampler
 - Relatively fast
 - Can be done from a boat
 - Samples collected in 50 cm sections as "undisturbed half cores"
 - Good horizon resolution and good for bulk density
 - Limited to soft materials
 - Small sample size
 - Moderate cost (\$1000)



Sampling

- Vibracoring
 - Excellent undisturbed cores
 - Up to several meters long
 - Can be used in dense materials
 - Can be stored for later examination
 - Slow and cumbersome (set up)
 - Costly equipment
 - □ Some "collapse" change in volume
 - Especially with organic rich horizons







Problems with Collapse

- Poor estimate of volume for bulk density measurement
- Only estimates of horizon depth



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Morphology and Horizons

- Equivalent of "young alluvial soils"
- Commonly are stratified
- Often have only A, C, and transitional (AC) horizons.
- Recognition and importance of the "oxidized" horizon at the surface





Oxidized Horizon (A)

Forms as a function of bioturbation and diffusion of oxygen



(From Payne, MS Thesis)

(From Payne, MS Thesis)

SPI and Redox



Not static – changes over time and with seasons.

Morphology and Horizons

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- Commonly are stratified
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- Recognition and importance of the "oxidized" horizon at the surface
- Most horizons are depleted (g) or gleyed
 - Some may be "reduced matrices" especially fresh systems



Subaqueous Soil Profile Description

Date	Latitude	N 38	Longitude	W 75	Water Depth
Core #					Predicted Tide
Describers DMB					
Time	Depth Inside Core			Compaction (Depth Outside Core - Depth Inside Co	re)
	Depth Outs	ide Core		Sampling Method	

	-																	
Horizonation	Bound	ary	USDA Text	ure	Matrix	Redox	imorphic Feat.	Organic Fragments		Stru	Structure			N-Value	Roots	Worm	Shells	Sulfur
	Depth	Dist	% gravel	Class	Color	Color	Abund/Contr	Abund	Color	Gr	Shape	Consis	t		%	Tubes	%	Odor

Morphology and Horizons

- Sometimes old (or truncated) soils can be buried with a younger Holocene age subaqeous soil.
- Submerged mainland slope landform

Preliminary Digital Elevation Model for Rehoboth Bay Bathymetry August 2003







Relict redoximorphic features from truncated, buried paleosol



Copppock, RB14







Copppock, RB14

Physical Properties

Particle size – Texture

n-Value

Bulk Density

Particle Size

Lab analyses, may need to first remove salts

- Chloride salts from sea water
- Sulfate salts generated during drying from oxidation of sulfide minerals
- Field texturing challenging because they are always too wet!





% Sand

From Balduff (2007)



Comparison of Field and Lab Textures Balduff (2007)

	Field Textures		Textures Based on Particle-Size Analysis										
n	(S	fS	LS	LfS	SL	fSL	vfSL	L	SiL	CL	SiCL	SiC
							%						
15	S		87		13	\setminus /							
9	fS	\mathbb{N}	100	\setminus /		\backslash							
16	LS	IV	38		56	\mathbf{V}	6						
10	LfS		70	\square	20	\land	10						
26	SL		4		35	31	19		12				
7	fSL			/	43	$\langle \rangle$	43				14		
1	SC			100									
18	L				11	11	28	6	33		11		
6	SiL					17					50		33
3	CL					33			33		33		
16	SiCL								13	18	19	50	
49	SiC				2	2			14	8	12	45	16
12	С					8		8	58	8	17		

Comparison of Field and Lab Textures Balduff (2007)

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n		S	fS	LS	LfS	SL	fSL	vfSL	L	SiL	CL	SiCL	SiC
							%						
15	S		87		13								
9	fS		100										
16	LS		38		56		6						
10	LfS		70		20		10						
26	SL		4		35	31	19		12				
7	fSL				43		43				14		
1	SC			100									
18	L				11	11	28	б	33		11		
6	SiL					17					50		33
3	CL					33			33		33		
16	SiCI								13	18	19	50	
49	SiC				2	2			14	8	12	45	16
12	C					8		8	58	8	17		



Distribution of particle-size data for 188 subaqueous soil horizons analyzed in Chincoteague Bay.



Distribution of particle-size data for subaerial soils found throughout Maryland (University of Maryland Pedology Lab, 2007). Each marker represents a different county in Maryland.

From Balduff (2007)

n-Value

From Soil Taxonomy (2nd Edition)

- The *n* value (Pons and Zonneveld, 1965) characterizes the relation between the percentage of water in a soil under field conditions and its percentages of inorganic clay and humus. The *n* value is helpful in predicting whether a soil can be grazed by livestock or can support other loads and in predicting what degree of subsidence would occur after drainage.
- For <u>mineral soil materials</u> that are not thixotropic, the *n* value can be calculated by the following formula:

$$n = (A - 0.2R)/(L + 3H)$$

- A = % water (field condition, on a dry-soil basis)
- R = % silt + sand
- L = % clay
- H = %OM

n-Value

From Soil Taxonomy (2nd Edition)

- but the critical *n* value of 0.7 can be approximated closely in the field by a simple test of squeezing a soil sample in the hand.
- If the soil <u>flows between the fingers with difficulty</u>, the *n* value is between 0.7 and 1.0 (slightly fluid manner of failure class); ...
- In the soil flows easily between the fingers, the *n* value is 1 or more (moderately fluid or very fluid manner of failure class).
- Soils in which the moisture content is periodically reduced below field capacity seldom have an *n* value of 0.7 or more. Most of the soils that have been permanently saturated are likely to have a high *n* value. Consequently, high *n* values are primarily in soils of tidal marshes, swamps, and shallow lakes. The sediments in these areas have never been above the capillary fringe during drought cycles.

... and subaqueous soils!

Cimg0082.mp4







Field Estimated n value

From Balduff (2007)



From Balduff (2007)

n-Value

- A useful parameter
 - Estimates bearing capacity
 - Implications for sedimentary environment
- Should be estimated in the field
- Little correspondence between field and lab values
- Lab values have little meaning, while field estimations are useful
- Rather than putting numerical values, should perhaps go with classes from SSM

From the Soil Survey Manual

Table 3-19. Manner of failure classes

Classes	Operation	Test Description Characteristics
* Deformable	Same	Specimen can be compressed to half its original thick- ness without rupture. Radial cracks may appear and extend inward less than half the radius normal compression.
Nonfluid	A handful of soil material is squeezed in the hand.	None flows through the fingers after exerting full compression.
Slightly fluid	Same	After exerting full compression, some flows through the fingers, but most remains in the palm of the hand.
* Moderately fluid	Same	After exerting full pressure, most flows through the fingers; a small residue remains in the palm of the hand.
Very fluid [*]	Same	Under very gentle pressure most flows through the fingers like a slightly viscous fluid; very little or no residue remains.
* The approximate Deformable Slightly fluid Moderately Very fluid-	e equivalent n-values, Pons a < < 0.7 n-value 1 0.7-1 fluid 1-2 ≥ 2	nd Zonneveld (1965), are as follows:

Bulk Density

Pretty convenient and easy if sampling with

- Vibracorer or MacCauley Sampler
- Cut linear section of core (or half core) and calculate volume.
- Dry and weight the sample



Bulk Density Pretty impossible if sampling with bucket auger



Chemical Properties

Sample Handling and Storage

- Physical properties pretty stable
- Chemical properties highly labile can change radically!



Sample Storage

- Bare minimum
 - Store on ice in a cooler
 - Refrigerate and analyze immediately
- Prefered
 - Sparge with N2
 - Then store on ice (or preferably) freeze in the field
 - Store frozen until analyzed
- Slows chemical and microbial oxidation of sulfide minerals



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Pedon	CB21
Landform	Submerged wave-cut headland
Soil Map Unit	Spβ
Current Soil Classification	Fine-Ioamy, Thapto-Histic Sulfaquents
(Proposed Soil Classification)	(Fine-Ioamy, Thapto-Histic Sulfiwassents)
Series	Southpoint Taxajunct
173 cm	C 260 cm

Black monosulfides (FeS) can form quickly by chemical reactions between Fe(II) and S=



Well formed and crystalline pyrite framboid

Pooly formed and poorly crystalline framboid formed within 2 years

Pyrite forms as a result of microbial processes, usually over periods of weeks, months, years or decades.

Oxidation of sulfides

FeS oxidizes very quickly - chemically



FeS₂ takes more time to oxidize – microbially mediated reactions.

Oxidation of sulfides

Overall oxidation and hydrolysis of S and Fe of pyrite

 $FeS_2 + 3 \frac{3}{4}O_2 + 2 \frac{1}{2}H_2O \rightarrow FeOOH + 2H_2SO_4$

- Sulfuric acid drives down pH
- Acid reacts with other basic compounds generating salts, so salinity goes way up.
- This is why proper handling and storage are so important!

Field Measurements

- pH mostly near neutral
 - can measure with standard field kits (Truog; test strips)
- FeS (acid volatile sulfide) qualitative
 - Check sample for H_2S odor
 - Add a few drops of 10% HCI and see if you detect increase in H₂S
 - If breezy, place small amount of sample in ziplock bag and add a few drops of HCI. Seal and wait a minute. Open and see if you detect increase in H₂S odor inside the bag.
 - Some experimentation with H_2O_2

Lab Measurements

- Moist, oxidizing, pH Incubation over time
 - Optimizes conditions for microbial oxidation of sulfide minerals and generation of acidity and salts



From Balduff (2007)

pH Incubation – Some samples take longer for pH to drop.

Length of time to drop	Number of	% of Samples that eventually					
below pH 4	samples⁺	show a drop in pH<4					
4 weeks	25 (16%)	20					
8 weeks	73 (45%)	57					
12 weeks	104 (64%)	81					
16 weeks	117 (72%)	91					
20 weeks	125 (77%)	98					
24 weeks	128 (78%)	100					
† 46 samples were only incubated for 16 weeks							

From Balduff (2007)

Lab Measurements

Salinity

- Electrical conductivity run on 1:5 extract
- Much simpler than saturated paste

(From Payne, MS Thesis)



Lab Measurements

- Sulfides Cr reduction methods
 - Several variations; fairly involved
 - Sample handling critical prior to analysis
 - Can distinguish between FeS (Acid Volatile Sulfide - AVS) and Pyrite (Cr reducible Sulfide -CRS) if needed



Usually, AVS (FeS) is MUCH LESS than CRS (FeS₂), typically (only a few % of CRS).

From Balduff (2007)



Carbonates – Qualitative

Add drop of 10% HCI to soil while examining under stereo microscope

Class	Reaction	CaCO ₃ (g kg ⁻¹)			
		Mean	Range		
non-effervescent (NE)	no reaction	0			
very slightly effervescent (VS)	one or two bubbles	3.2	0.0 to 17		
slightly effervescent (SL)	few bubbles	7.4	0.0 to 30		
strongly effervescent (ST)	many bubbles		18 to 370		
violently effervescent (VE)	low foam				

From Balduff (2007)

